## DEPARTMENT OF EARTH AND ENVIRONMENTAL SCIENCES

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#### Abstract

The major phase of the Earth's mantle, (Al,Fe)-bearing bridgmanite, transitions to the post-perovskite (PPv) phase at Earth's deep lower mantle conditions. Despite extensive experimental and ab initio investigations, there are still important aspects of this transformation that need clarification. Here, we address this transition in  $(AI^{3+}, Fe^{3+})$ -,  $(AI^{3+})$ -,  $(Fe^{2+})$ - and (Fe<sup>3+</sup>)-bearing bridgmanite using ab initio calculations. We find that the seismic features produced by the PPv transition depend distinctly on the chemical composition. For instance,  $Fe^{3+}$ -,  $Al^{3+}$ -, or ( $Al^{3+}$ ,  $Fe^{3+}$ )-alloying increase the transition pressure, while  $Fe^{2+}$ -alloying has the opposite effect. Consequently, the absence of a D" seismic discontinuity or signature of a double-crossing of the PPv phase boundary point to a Fe<sup>2+</sup>-poor and Fe<sup>3+</sup>-rich bridgmanite composition with profound implications for the redox state of the deep lower mantle. These chemistry-specific seismic features together, along with thermochemical equilibrium calculations will be fundamental for resolving the chemical composition of the D" region by direct inspection of tomographic images.

### Quasi-ideal solid solution model

A: MgSiO<sub>3</sub>

B:  $Mg_{0.875}Fe^{2+}_{0.125}SiO_3$ 

(Mg<sub>0.875</sub>Fe<sup>3+</sup><sub>0.125</sub>)(Si<sub>0.875</sub>Fe<sup>3+</sup><sub>0.125</sub>)O<sub>3</sub>

 $(Mg_{0.875}AI_{0.125})(Si_{0.875}AI_{0.125})O_3$ 

(Mg<sub>0.875</sub>Fe<sup>3+</sup><sub>0.125</sub>)(Si<sub>0.875</sub>Al<sub>0.125</sub>)O<sub>3</sub>

Gibbs free energy of mixing curves of the Pv and PPv states in a A-B binary system:

$$\Delta G_{M}^{Pv} = k_{B}T(X_{B}\ln[X_{B}] + (1 - X_{B})\ln[1 - X_{B}]) + X_{B}(\Delta G_{B}^{Pv - PPv})$$

$$\Delta G_M^{\rm PPv} = k_B T (X_B \ln[X_B] + (1 - X_B) \ln[1 - X_B]) + (1 - X_B) (\Delta G_A^{\rm Pv-PF})$$

The compositions of the Pv and PPv solvus lines

$$X_B^{PPv} = \frac{1 - \exp\left(\frac{G_A^{PPv} - G_A^{Pv}}{k_B T}\right)}{\exp\left(\frac{G_B^{PPv} - G_B^{Pv}}{k_B T}\right) - \exp\left(\frac{G_A^{PPv} - G_A^{Pv}}{k_B T}\right)}$$
$$X_B^{Pv} = X_B^{PPv} \times \exp\left(\frac{G_B^{PPv} - G_B^{Pv}}{k_B T}\right)$$

The PPv fraction,  $n_{PPv}$ , is given by the lever rule

$$n_{\rm PPv} = \frac{X_B - X_B^{\rm Pv}}{X_B^{\rm PPv} - X_B^{\rm Pv}}$$

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Figure 4. PPv fraction  $n_{PPv}$  for all bearing cases, with a bearing element concentration x=0.10. The LDA boundaries are generally wider than the GGA boundaries. The differences between the two would result in phase boundary uncertainties.

# The post-perovskite transition in Fe- and Al-bearing bridgmanite

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### Methods and calculation details



Figure 2. Elasticitc properties [1,2,3]



Noted here  $X_B = \frac{x}{0.125}$ , and x is the alloying element concentration.

110 130 150 170 Pressure (GPa)

Figure 3. Schematic of the two Gibbs free energy of mixing curves of the Pv and PPv states in an A-B binary system.



- Fe<sup>2+</sup> reduces the transition pressure, while Al<sup>3+</sup>, Fe<sup>3+</sup>, (Fe<sup>3+</sup>, Al<sup>3+</sup>)- increase it. The coexistence between Pv and PPv phases grows in proportion to the bearing element concentration.
- The calculated changes in acoustic can serve as a guide to unravel details about the composition of the deep lower mantle.

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- Light grey shaded areas are the uncertainties from LDA and GGA data.
- CO9 Ref.[5]; T05 Ref.[6]; Geotherm<sup>1</sup> Ref.[7]; Geotherm<sup>2</sup> Ref[8].
- The dark colors are the most likely areas of coexsistence between the pv and ppv phases
- The ligher colors are the uncertainties brought by the DFT calculations.
- A10 Ref.[9]; T19 Ref.[10]; C08 Ref.[11]; M12 Ref.[12]

### Conclusions

#### Depth and thickness of Pv-PPv boundaries Experiments



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